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ARTICLE Internal Waves Formation and Propagation in the Persian Gulf

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ABSTRACT

The Persian Gulf (PG), as a semi-enclosed water basin extends in [47-57] E, [24-30] N, geographic domain. Particularly, northern part of the PG shows more baroclinicity and turbulence because of the river inflow from the Arvand, bottom and costal stresses. Furthermore, wind stress has many effects rather than in mid deep domain of the PG. Thermocline development in the PG is observed because of studying the data measured in the Mt. Mitchell cruise in 1992 by different models from winter to summer. The studied turbulence in the northern part of the PG is navigated from winter to summer due to the internal wave's activity and stability intensified through water column.

1. Introduction

Internal waves play a major role in the formation of seawater stratification and are responsible for the main processes of ocean dynamics, such as energy transfer and mixing ^[8]. The surface layer is influenced by local wind forcing. The thermocline tries to re-stabilize itself after the cessation of dry season and oscillates for the rest of the year. A main source of turbulence energy is the internal wave field. Some of internal wave energy derives from the transmission of wind energy through the surface mixed layer ^[1]. The measurement stations in the PG are presented in Figure 1.





(Northwestern, mid deep and near the Strait of Hurmoz)

An important mechanism for exiting fields of internal

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gravity waves (IGW) in natural (ocean, the Earth atmosphere) and artificial stratified media is their generation by sources of perturbation of various physical nature, i.e., of natural (moving typhoon, wind waves, flow past the ocean bottom relief imperfections, variations in the density and flow fields, leeward mountains) and anthropogenic) marine technological structures, collapse of the turbulent mixing region, underwater explosions(character ^[6]. As a result of variations of physical characteristics of PG water through column, and effects of climate on the water surface, internal waves happen and vary in water column of PG. Therefore, study of the waves in the PG water column has an importance we concentrate in this paper including results of a research.

2. Thermocline Formation in the PG

Thermocline forms in the PG, and this is a seasonal phenomenon because of particularly temperature variation through water column in the PG. Therefore, thermocline will be studied in this research according to data collected during Mt. Mitchel cruise in the PG, curves contours are drawn in Figure 2 to Figure 6. Temperature, density and salinity are three useful physical properties of seawater to studying. Naturally, the temperature of seawater is fixe at the surface by heat exchange with the atmosphere. Temperature and density vertical variations in some stations in three measurement sections are as in the following profiles in the PG in winter and summer.



Figure 2. Vertical variations of temperature and density in three measurement sections in the PG, winter and summer 1992

The atmospheric disturbance is set by an idealized pulse of cyclonic wind stress with a Rankine vortex structure. Strength, radius and duration of the forcing are varied. The effect upon wave generation of stratification with variable mixed-layer depth is also examined ^[9]. Relying on the accuracy of the curves in Figure 2, it could be stated that vertical gradient of temperature and density in the subsurface of the water column is very high in the summer but in winter, it is high only in the Strait of Hormuz because of water exchange in winter.

In fact, it could be resulted that thermocline exists only in summer. Overall, the temperature breaks in the subsurface layer through the water column because of wind stress, demonstrates the thermocline developed in the study area. In this regard, thermocline formation and its development are associated to some variations in internal waves, particularly in subsurface layer.

In fact, thermocline development and its variation in space and time represent a turbulent flow, and thus propagation of internal waves through water environment. Turbulence and internal waves propagate along the baroclinic pressure gradient in a vertical direction through the water column due to baroclinicity and tidal wind stress ^[10]. Based on Figure 2 comparing the curves, we could observe that thermocline does not form in winter, particularly in the Strait of Hormuz because of water exchange with the Gulf of Oman. However, it forms all over the PG through the water column, suggesting that thermocline develops from the Strait of Hormuz to the PG North from winter to summer.

3. Results and Discussion

The point is that internal waves can create rather strong currents on the ocean surface. The flow is changed depending on the wave extension: its velocity is greater at the wave crest and wave trough, and is slower where thermocline oscillations are little ^[7]. Internal waves in PG in summer differ from winter, due to differences of forcing and factors in summer and winter. Internal waves breaking results in turbulence propagation during winter to summer ^[2]. Arvand river inflow in northwestern of the PG helps to stratification, and then internal waves and turbulence can form there. Of course, the gradient throughout density contours throughout the PG represents the existence of the internal waves due to thermocline formation, particularly in summer because of high intensity forcing. Wind stress affects the PG surface in the following processes:

(1) An atmospheric disturbance causes an internal wave.

(2) Turbulence is driven by internal waves shear.

(3) Vertical density fluxes are resulted from the mixing.

(4) A gravitational imbalance in the horizontal direction includes an intrusion.

(5) High wave number internal waves are generated at the head of the intrusion.

Wind stress as a forcing generates an acceleration for water body in vertical and horizontal directions with its components due to shear induced on water ^[5].

Thermocline forms in the open ocean and can be a seasonal phenomenon in the shallower part of the ocean or a permanent since it is seen in the deeper part of the open ocean. Thermocline is a phenomenon which happens in the Persian Gulf., and it develops from winter to summer resulted of the numerical model (POM) in the PG^[3]. Formation of a shallow thermocline with upper mixing and displacements of the thermocline is associated with internal wave propagation, which plays an important role in turbulence generation. The stratification as an important factor influencing on local generation of internal waves in the coastal zones subjected to seasonal variability ^[4].

In this study, we have wanted to appoint internal waves dependence on thermocline development in the PG, and so will conclude internal waves and turbulence. As in the above profiles, there are fluctuations and variations along vertical T & D; therefore, it could be a crucial result of internal wave's existence. Thus, thermocline development often causes turbulence. We could express that turbulence happens in the PG, because of some forcing all over the PG, such as: tide, wind stress, river inflow. In the PG, baroclinic pressure happens in various stations as in the last profiles. Buoyancy frequency (Brunt Vaisala) as following as baroclinicity and density gradients received by a code in FORTRAN in some stations in regards to turbulence and internal wave's existence are drawn against depth. Figure 3 shows vertical variations of buoyancy frequency in two measurement sections. In fact, in most of the measurement stations, internal waves could form.



Figure 3. Buoyancy frequency N (z) (* 0.01) in some important stations in summer and winter 1993, winter and summer 1992

In fact, more variations through the buoyancy frequency happen in summer rather than in winter, because of more gradients in T and D. The buoyancy frequency is calculated as in the following:

$$N(z)^{2} = (-g/\rho) (\Delta \rho / \Delta z)$$
(1)

Where w, N (z), w and p (z) are vertical component of internal wave velocity, wave number velocity and mean density function in water column. Depth mean variations for internal wave's velocity (vertical component) profiles in the PG is received as in Figures 4 and 5 in summer and winter with and without bottom stress.



Figure 4. Mean TKE through the PG during winter until summer



Figure 5. Mean TKE through the PG during winter until summer without bottom stress, winter and summer 1992

In figure 6, profiles of vertical component of internal wave velocity in some points of the PG in winter and summer 1993 depth (m) against velocity (m/s) are showing turbulence in summer because of thermocline develops from winter.



Figure 6. Internal wave's velocity variations in some stations in the Persian Gulf in winter and summer, winter and summer 1992

As it could be seen, internal wave's velocity variations in z direction are more in summer rather than in winter, because of thermocline fluctuations based on much density variations.

4. Conclusion

Existence of stratification and development of thermocline in the PG is a crucial phenomenon. Summer thermocline formation in the PG is a seasonal potential of turbulence, and internal wave's creation would happen in the case study zone. Thermocline formation is usually associated with internal waves as an important cause in turbulence.

Internal waves are able to provide and transmit the energy, which is essential for aquatic circulation and mixture. Moreover, in coastal areas that charges in fluid temperature in water column causes diffusion of internal waves.

It could be concluded that:

(1) Baroclinic pressure and stratification, cause turbulence throughout PG water column in summer.

(2) In summer, turbulent behavior happens. The most effective factors on the behavior in the PG are tide, wind, evaporation and river inflow.

(3) The existence of tides, northwestern winds and strong internal waves in PG results in high vertical component of internal wave's velocity.

(4) Turbulence and internal waves due to baroclinic pressure because of density gradients exist in the PG water in summer.

(5) Vertical component of internal waves velocity has range of [-0.8, +0.8] with more fluctuations in summer.

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